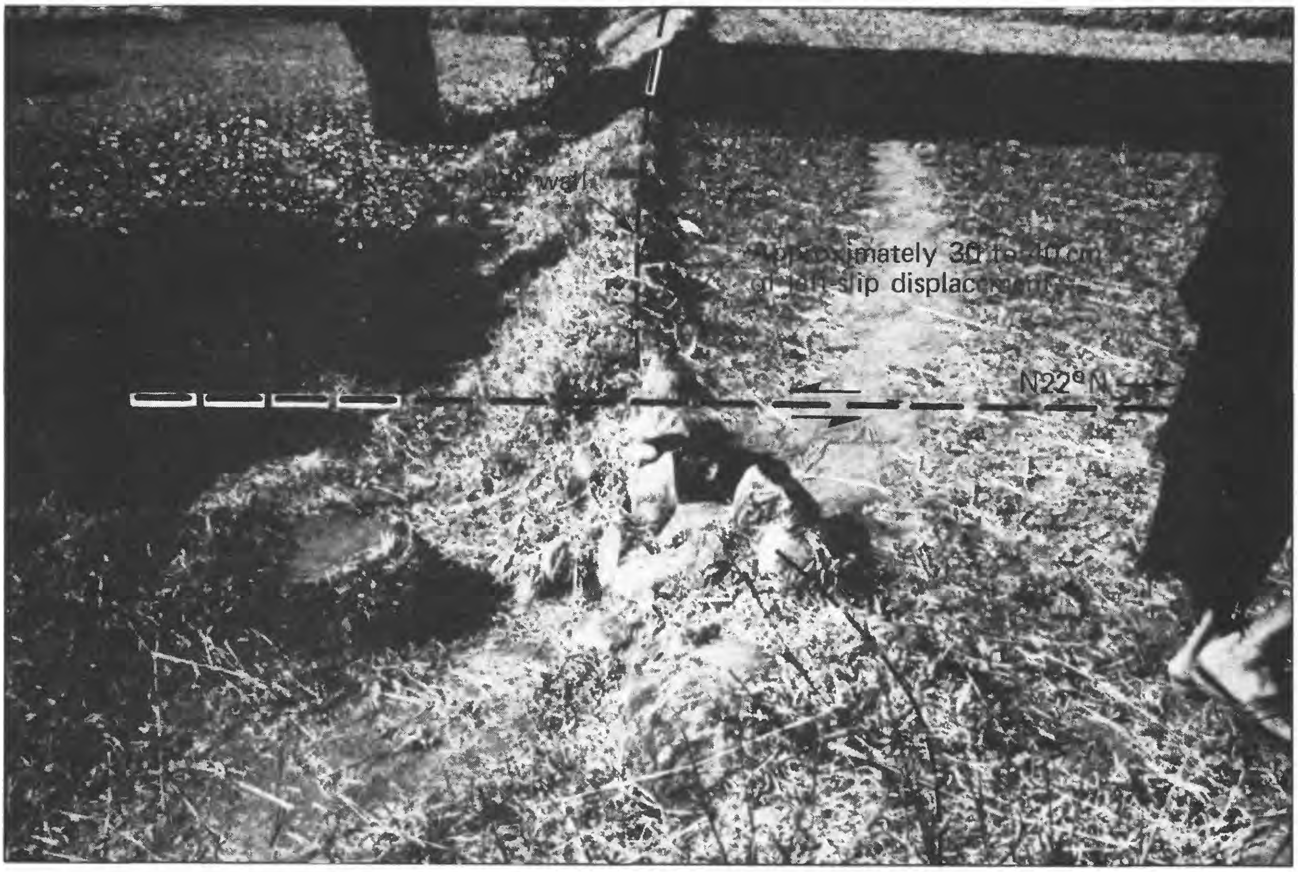


OFFSET PADDY WALLS ALONG
ALIGNMENT OF GROUND CRACKING

India

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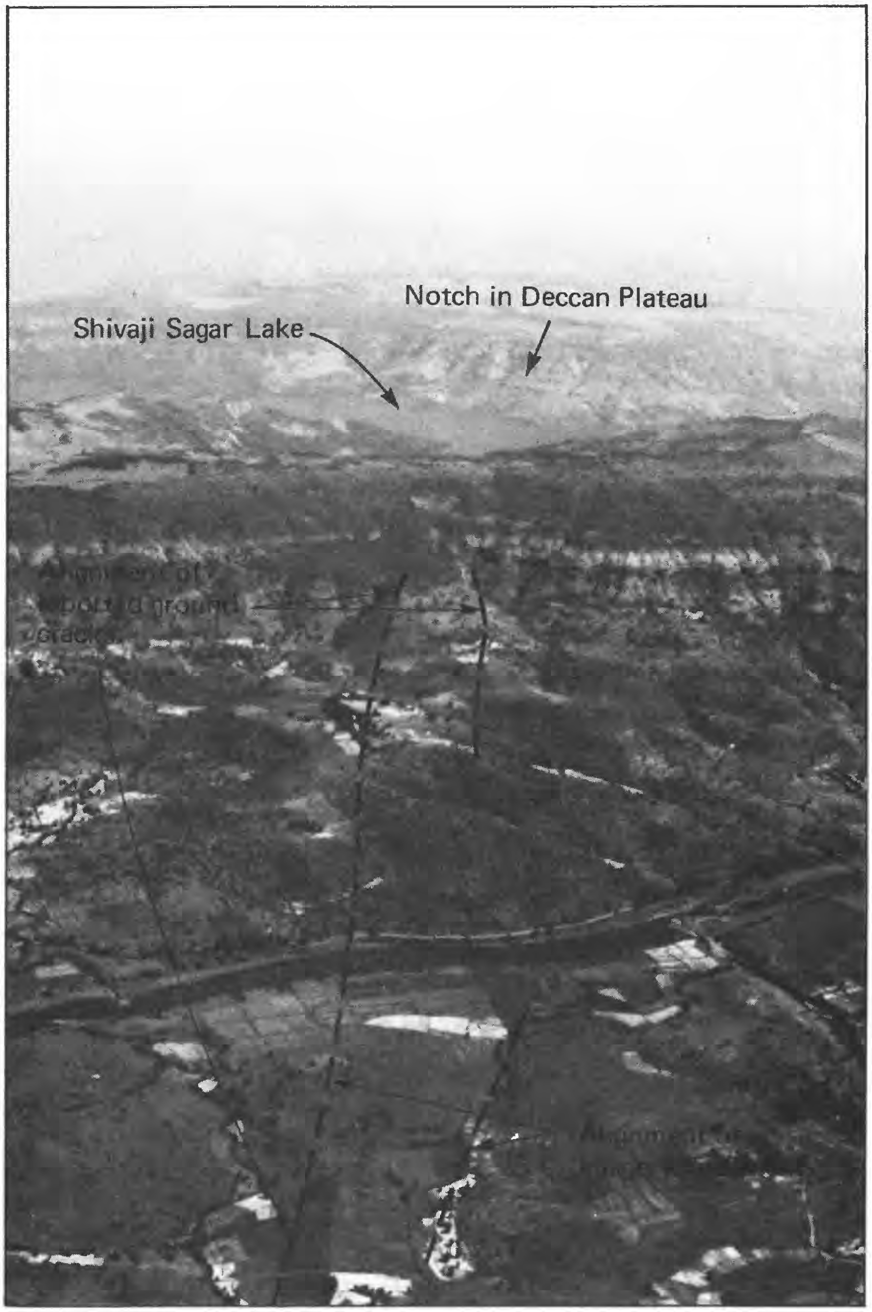
| Project 14087A | Figure 4-5 |
| **Woodward-Clyde Consultants** | Page 122 |




TOPOGRAPHIC NOTCH IN DECCAN PLATEAU
ALONG ALIGNMENT OF GROUND CRACKS
VICINITY OF KOYNA DAM
India

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| Project No. 14087A | Figure 4-7 |
| **Woodward-Clyde Consultants** | Page 129 |

cracks at the location indicated by the natives. Analysis of data showed the natives had pointed out a trend of ground cracking to the west of the zone noted in 1977. The two zones of cracking observed high on the hillside north of Donechiwadi correspond to the two separate trends of ground cracking pointed out by the natives as passing east of Donechiwadi and passing through Donechiwadi.

Approximately 500 m south of the village of Donechiwadi, the reported remaining surficial expression of a ground crack was exposed. This feature is oriented normal to the slope of the low hill at the site and is 3 to 4 m long, 30 cm wide, and 15 cm deep. The villagers stated that the feature was a ground crack related to the earthquake along which erosion subsequently has occurred. The orientation of the feature relative to the slope direction does not preclude the feature being an erosional gully. Examination of the feature did not resolve the nature of its origin.

Ground cracking along the road between Koyna and Karad, immediately adjacent to the "48 km to Karad" road marker (Figure 4-4), has been reported previously (Committee of Experts, 1968). Approximately 100 m east of the road marker and the zone of ground cracking is an outcrop on the north side of the road [see also Cluff (1977) description of the location]. Stories conflicted regarding the location of the zone of ground cracking relative to this outcrop. Initially the zone was shown to cross the road at the east end of the outcrop. Subsequently, the zone is shown to cross at the west end of the outcrop, on the west facing slope of the outcrop and the adjacent gully. On the basis of this investigation, the second location is judged to be the more accurate because many, if not all, of the villagers present tended to overrule the eastern location, although the eastern location may represent the easternmost zone of the two trends of ground cracking.

The outcrop was examined in detail and consisted of prominently jointed basalt. No evidence of fault displacement was observed. The predominant joint orientation was N15°W to N20°W and had generally a steep northeast dip. Joints were observed to be generally tight and devoid of infilled or secondary material. In addition, no fault gouge, slickensides, or shear zones were observed to be present. The surficial soil units were thin to nonexistent. Erosion appears to be stripping away any soil cover that has developed.

Whether the origin of the western zone of ground cracking in the vicinity of Donechiwadi was from differential settlement, slumping, or fault displacement could not be readily determined from this examination. To assess if the ground cracks were of tectonic origin, a location would have to be selected for examining the zone for evidence of active faulting; finding such a location in a short period of time in the area around Donechiwadi was considered unlikely. Consequently, the zone of ground cracking (reported by the Committee of Experts, 1968) was examined for such evidence on the south side of the Koyna River near the village of Kadoli.

<u>Kadoli</u> - The villagers indicated the location of the zone of ground cracks in the vicinity of Kadoli (Figure 4-4, south of the Koyna River). This location appears to contradict Plate 5-1 in the report of the Committee of Experts (1968), unless the arrow (on the plate) showing the direction to Kadoli is reversed 180 degrees. The villagers were very emphatic in locating the ground cracks at the location shown on Figure 4-4, and the described zone was oriented N30°E. South of the Rain Temple, the ground cracks were observed to have continued for approximately 500 m to the crest of the hills south of the temple. The villagers did not know if the ground cracks continued south of the hill crest because no

villager had gone to the other side of the hill to look for them. The zone was reported to have been 10 m wide and to have consisted of a nearly continuous central crack flanked on either side by one to three additional cracks. When requested to draw what was observed, the villagers drew a configuration with no apparent en echelon stepping of the cracks. The main crack was reported to have been 10 cm wide and 70 to 100 cm deep. The flanking fractures were 2 to 5 cm wide, 70 to 100 cm deep, and 2 to 3 m long.

Between the Rain Temple and the Koyna River, the zone of ground cracks crossed two stream drainages. The southern-most drainage and the Koyna River bank both were examined for evidence of faulting, but none was observed. In the northern drainage, however, a fault was observed within or adjacent to the reported zone of ground cracking (Figure 4-6). The fault was located in the northeast bank of a small stream drainage south of the Koyna River (Figure 4-8). Its orientation was N35°E and nearly vertical. Close inspection of the fault showed two distinct near-vertical shears in weathered basalt. The upper shear has a strike of N48°E and dips 79°NW. The lower shear strikes N35°E and dips 83°N. Both shear zones contain zones of clay up to 8 cm thick. The surface of the margin of the clay zone and other planes within the clay are polished. Slickensides were observed on some planes and on the edge of a small cobble adjacent to and in contact with the clay zone. A polished pebble was found in the clay zone with the long direction of the pebble oriented parallel to the grooves in the slickensided clay planes (Figure 4-9). Well-defined manganese-stained slickensided planes were measured to have a strike of 35°E and a rake of 31°SW. The clay-filled zone was traceable for 1.6 m from the base of the stream bank upward through weathered basalt to near the ground surface where slumped reddish brown soil (pebbly colluvium) obscured the structure, as shown in Figures 4-8 and 4-9.


FAULT PLANE EXPOSED IN
BEDROCK ALONG RIVER BANK
VICINITY OF KOYNA DAM
India

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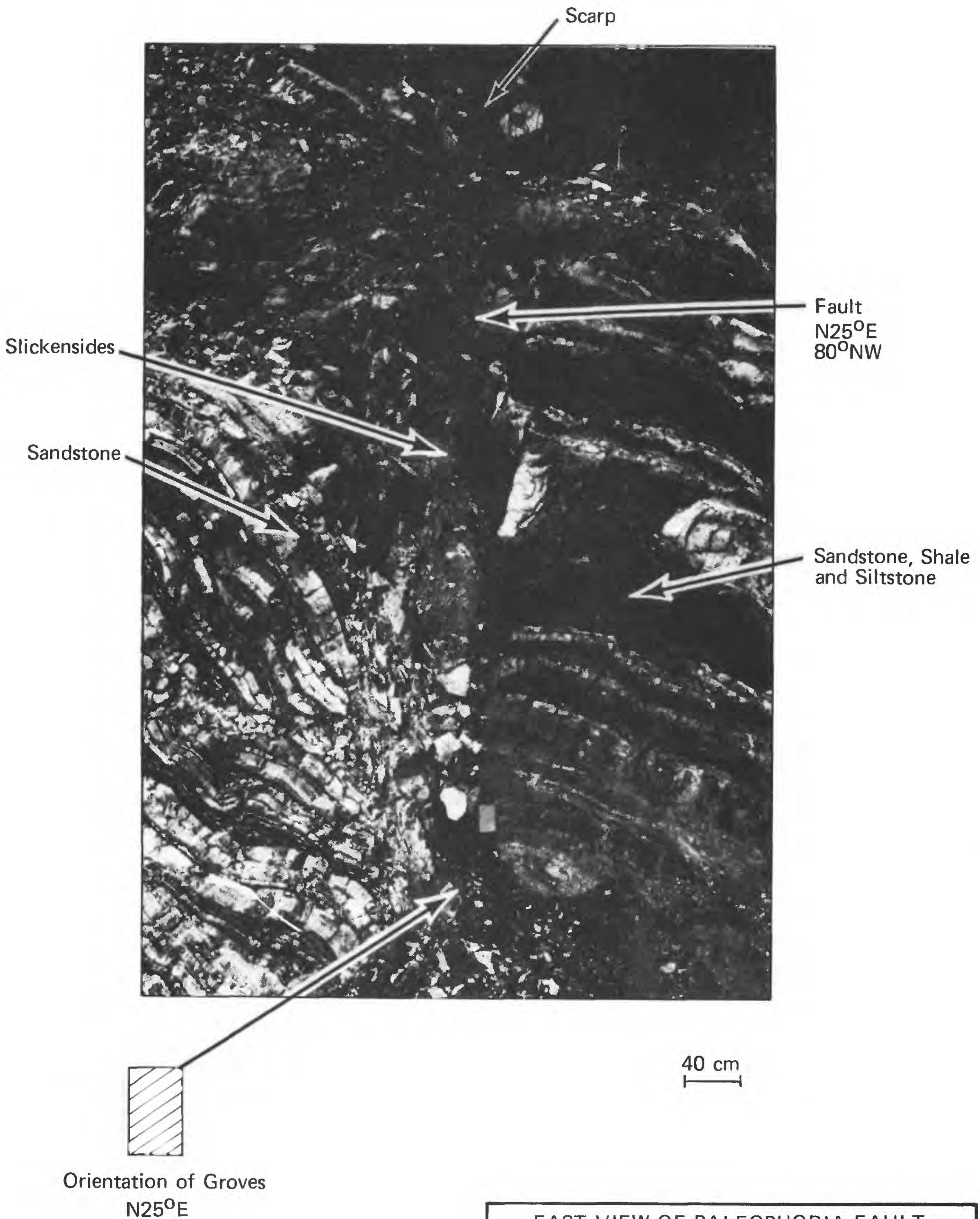
| Project No. 14087A | Figure 4-8 |
| **Woodward-Clyde Consultants** | Page 133 |


SLICKENSIDES ON FAULT PLANE
VICINITY OF KOYNA DAM
India

Project No. 14087A
**Woodward-Clyde Consultants**

Figure 4-9




EAST VIEW OF PALEOPHORIA FAULT
NEAR MEGDHOYAS BRIDGE
VICINITY OF KREMASTA-KASTRAKI
Greece

| | |
|---|---|
| Project No. 14087A | Figure 4-16 |
| **Woodward-Clyde Consultants** | Page 153 |


View is to the northwest

AERIAL VIEW OF ALEVRADA–
SMARDACHA FAULT THROUGH
THE VILLAGE OF ALEVRADA
Greece

Project No. 14087A
**Woodward-Clyde Consultants**

Figure 4-21

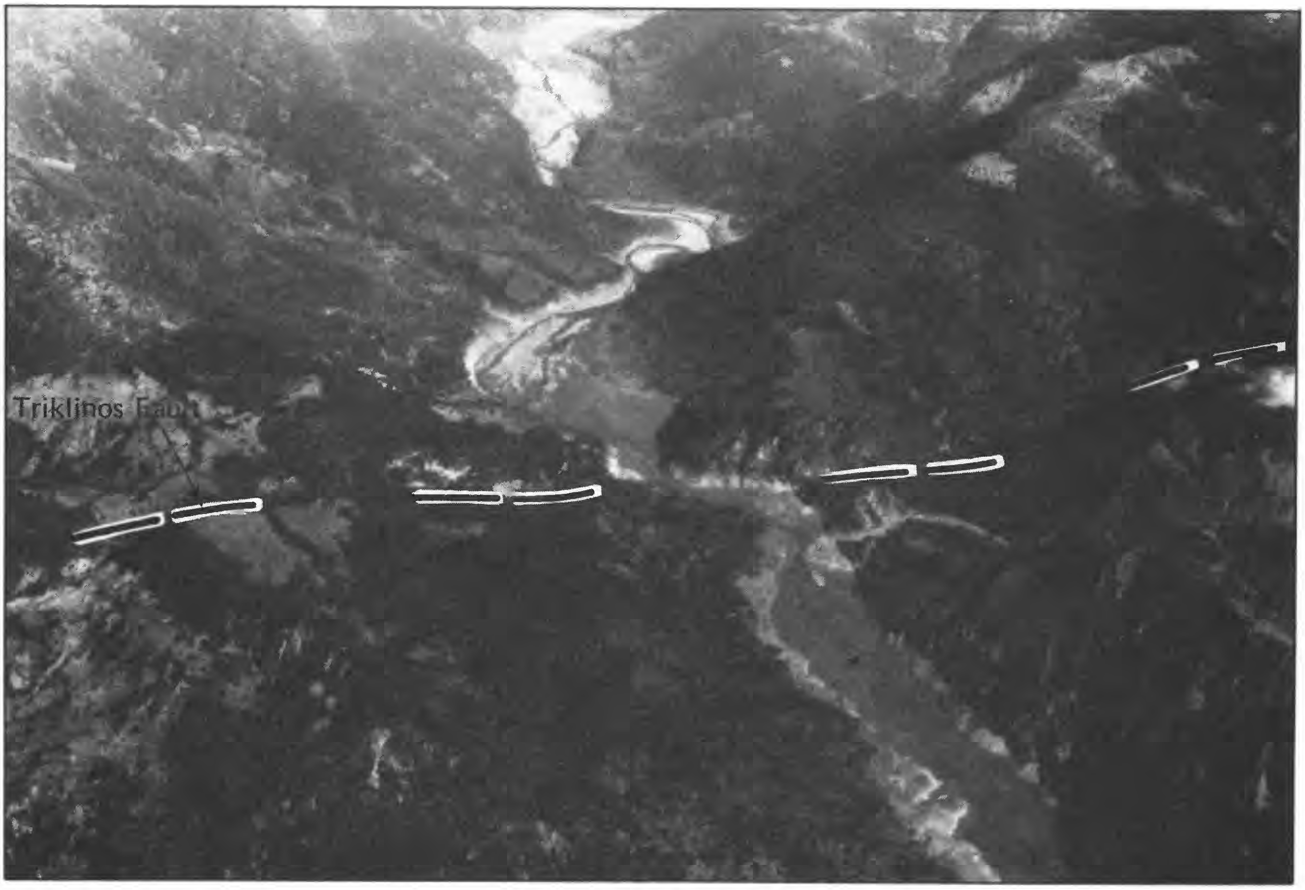
was observed to trend from the Megdhoyas arm of the reservoir, through the stream drainage from Frangista, which has had a sinistral displacement of approximately a hundred meters (Location D, Figure 4-14; Figure 4-22). The linear feature continues northwestward near the prominent scarp shown on Figure 4-15 and subparallel to the trend of the Pindus fault in this area of the lake region. Northwest of the village of Paleohoria, the Pindus thrust fault trends more northerly while the linear feature continues across the Acheloos arm of the reservoir and offsets the Alevrada-Smardacha fault with sinistral displacement (Location E, Figure 4-14; Figure 4-23). The linear trend was observed to continue northwestward to the village of Triklinos beyond which it could not be observed.

The Triklinos fault's observed length is at least 25 km. The fault trends northwest and is inferred to be active, based on the youthful nature of the displaced Frangista drainage. Recurring displacement within the present tectonic stress regime has been inferred to have occurred, resulting in the offset drainage of the stream. The trend of the fault and the sense of displacment agree remarkably well with that postulated by Comninakis and others (1968) (Figure 4-10). Correlation of the field observations and the seismologic calculations is not possible at the present time, but the similarity is notable.

SUMMARY AND CONCLUSIONS

1) Faulting is prevalent in the Kremasta reservoir area.

2) The Pindus thrust fault displaces Miocene sedimentary strata and passes through the reservoir area. The time of latest displacement along this fault could not be determined.




View is to the north.

AERIAL VIEW OF SINISTRAL DISPLACEMENT OF FRANGISTA STREAM DRAINAGE AT TRIKLINOS FAULT
Greece

| | |
|---|---|
| Project No. 14087A<br>**Woodward-Clyde Consultants** | Figure 4-22 |
| |  |


View is to the west.

| AERIAL VIEW OF SINISTRAL DISPLACEMENT OF THE ALEVRADA-SMARDACHA FAULT BY THE TRIKLINOS FAULT<br>Greece | |
|---|---|
| Project No. 14087A | Figure 4-23 |
| **Woodward-Clyde Consultants** | Page 166 |

3) The Paleohoria fault is a northeast-trending, high-angle fault whose regional extent was not determined. Displacement is up to the northwest. Field data suggest the more recent movement may be dextral and reverse. The fault is located in the eastern part of Lake Kremasta, and its trace projects through the Megdhoyas arm of the lake. Circumstantial evidence, including a topographic scarp and slickensides, suggests that it may be active.

4) The Alevrada-Smardacha fault is a prominent northeast-southwest trending, high-angle fault of at least 15 km length between Petrona and the Smardacha Bridge. Normal displacement down to the southeast was observed, as was sinistral and dextral displacement. Extensive landsliding occurred north of the village of Alevrada during the 1966 earthquake. The landslides apparently were triggered by the earthquake. No ground cracks were identified along the Alevrada-Smardacha fault in the Alevrada area. Clear topographic features and recent seismicity suggest this fault may be active.

5) The Triklinos fault is a northwest-trending, high-angle fault whose observed length is at least 25 km from the Megdhoyas arm of Lake Kremasta to the village of Triklinos. An offset stream drainage and offset of the Alevrada-Smardacha fault show sinistral displacement of approximately a hundred meters. The youthful nature of the displaced stream demonstrates the recency of fault displacement along this feature and strongly suggests the fault is active.

6) The Triklinos fault is approximately parallel to, and very near, the fault on which the 1966 earthquake (M = 6.3) was postulated to have occurred by Comninakis and others (1968). The sense of displacement is also the same.

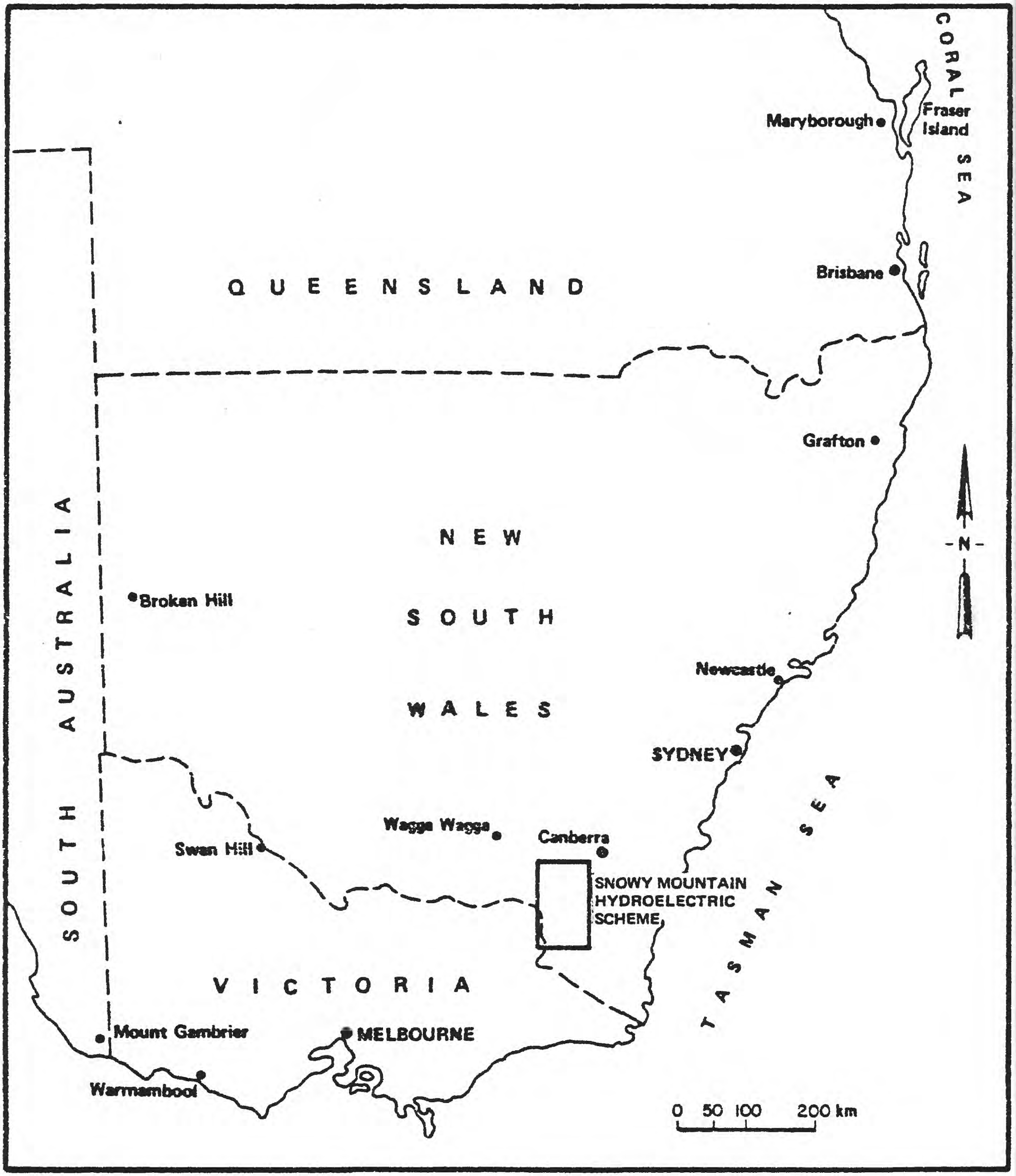
7) Active faults, including the Triklinos fault and possibly the Alevrada-Smardacha fault and the Paleohoria fault, are present within the hydrologic regime of Lake Kremasta.

## 4.4 EUCUMBENE, TALBINGO, AND BLOWERING RESERVOIRS, AUSTRALIA

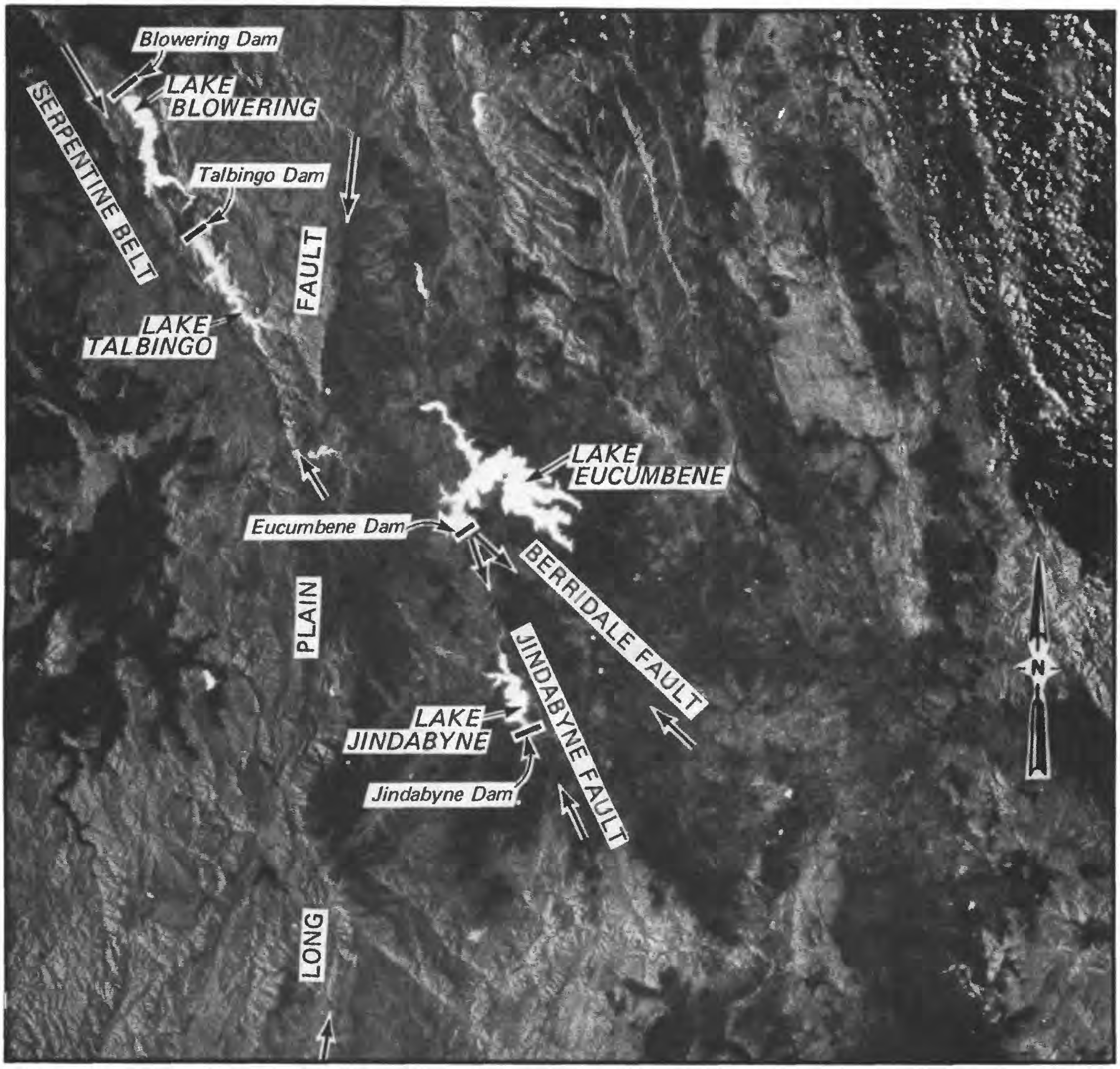
### INTRODUCTION

Lakes Eucumbene, Talbingo, and Blowering are the largest reservoirs in the Snowy Mountain Scheme, a hydroelectric and irrigation complex located in southeastern Australia (Figure 4-24). These lakes are formed by the impoundment of the Snowy and Eucumbene Rivers which are diverted through a series of tunnels to the Murray and Murrumbidgee Rivers (Figure 4-25). The characteristics of each reservoir are presented in Table 4-2, and a photograph of Eucumbene Dam is shown in Figure 4-26.

Initial impoundment of Lakes Eucumbene and Talbingo were accompanied by high levels of seismic activity that appear to be related to the reservoir filling. No unusual seismic activity was recorded during impoundment of Lake Blowering, which is immediately downstream from Lake Talbingo and which was impounded two years earlier than Lake Talbingo. A preliminary study was undertaken to evaluate fault activity within the regime of these reservoirs. For the purposes of this study, a fault is defined as active if it has had displacement during that area's present stress regime.




| PROJECT LOCATION MAP<br>LAKES EUCUMBENE, TALBINGO,<br>AND BLOWERING<br>New South Wales, Australia | |
|---|---|
| Project No. 14087A<br>Woodward-Clyde Consultants | Figure 4-24<br>Page 170 |




PROJECT AREA
LAKES EUCUMBENE, TALBINGO, AND BLOWERING
New South Wales, Australia

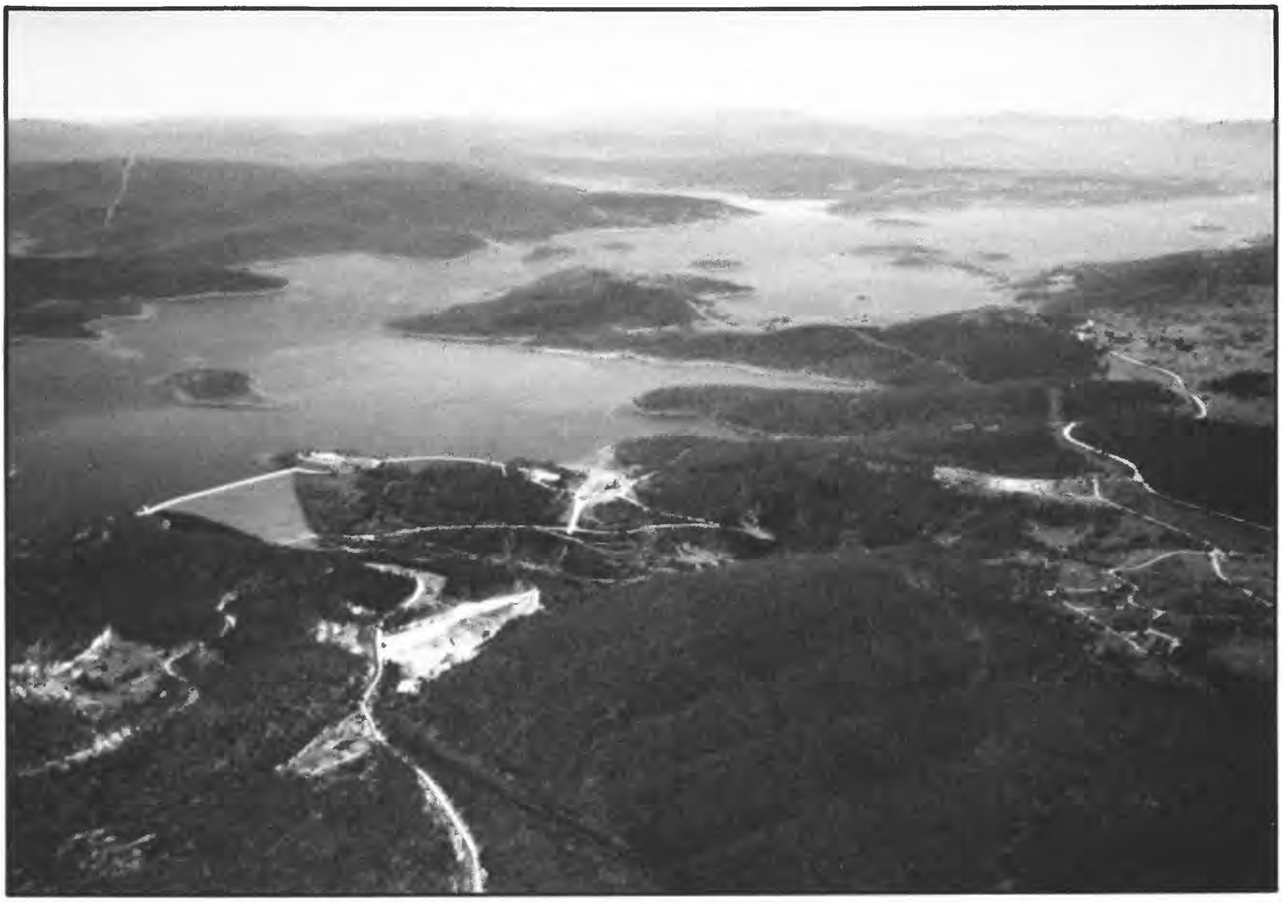
| Project No. 14087A<br>**Woodward-Clyde Consultants** | Figure 4-25<br>Page 171 |
|---|---|

TABLE 4-2
DESCRIPTION OF THE EUCUMBENE, TALBINGO, AND BLOWERING RESERVOIRS

| | Eucumbene | Talbingo | Blowering |
|---|---|---|---|
| Type | Earth/Rock | Earth/Rock | Earth/Rock |
| Water Depth | 106 m | 142 m | 95 m |
| Crest Length | 579 m | 700 m | 888 m |
| Reservoir Area | 145 $km^2$ | 19.4 $km^2$ | 44.5 $km^2$ |
| Gross Storage | 4761 x $10^6$ $m^3$ | 935 x $10^6$ $m^3$ | 1628 x $10^6$ $m^3$ |
| Date Completed | May 1958 | October 1970 | September 1968 |
| Foundation Rock | siltstone and quartzite | rhyolite lava and tuff | quartzite, metasiltstone, phyllite |

The scope of work for this study consisted of the following:

- gathering and review of published and unpublished data and maps;
- discussions with scientists resident in, or familiar with, the study areas;
- aerial reconnaissance flights over the study area;
- ground reconnaissance at selected locations in the study area;
- examination of aerial photographs of portions of the study area;
- collecting and evaluation of seismologic data.



View is to the northeast, across Lake Eucumbene. Eucumbene Dam is in foreground.

EUCUMBENE DAM AND
LAKE EUCUMBENE

New South Wales, Australia

| Project No. 14087A<br>**Woodward-Clyde Consultants** | Figure 4-26<br>Page 173 |
|---|---|

principally from the Snowy Mountains Hydro-Electric Authority and was based on data from a seismograph network established in 1957. Originally, this network consisted of four stations, each with a recorder, and was designed with the intent of ensuring detection and location of microearthquakes and larger seismic events. The data collected by the network are now telemetered directly to a recorder at the Australian National University in Canberra.

Lake Eucumbene

Impoundment of Lake Eucumbene began in June 1957, prior to establishment of the Snowy Mountains seismographic network. The pre-existing seismograph network was not capable of detecting events with magnitude less than 4 in the Snowy Mountains region. Thus, a good pre-impoundment data base is not available. However, a large body of seismic data has been accumulated since impoundment. Forty-four earthquakes of sufficient magnitude to be located accurately occurred during a 3-1/2 year period from 1958 to 1962 in a region located generally south to southeast of the reservoir (Figure 4-28). The largest of these events, which occurred near the community of Berridale in May 1959, had a magnitude of 5. Twenty-one minor shocks occurred in the vicinity of this magnitude 5 event, and they are reported to have the strain release pattern of a typical aftershock sequence. These shocks were relatively distant from the reservoir.

At the time of the main Berridale event, Lake Eucumbene had not yet reached the dead volume of approximately 432 x $10^6$ $m^3$; ponding of the live storage began in October 1958. Cleary and others (1964) interpret this as evidence that reservoir impoundment could not have been responsible. However, in comparison, Lake Mendocino in California, which has a total storage capacity of only 151 x $10^6$ $m^3$, is a well-documented probable source of reservoir induced seismicity (Toppozada and Cramer, 1978). Lake Marathon in Greece, which has a capacity of only 41 x $10^6$ $m^3$, is another documented probable source of reservoir induced seismicity (Gupta and Rastogi, 1976).

Fault-plane analysis of the Berridale earthquake suggests that the event originated on either a high-angle reverse fault with a strike of about N50°W degrees, or a low-angle, south- or southeast-dipping reverse fault with a possible range of strikes between N40°W degrees and east-west (Cleary and others, 1964). The first solution, which suggests a high-angle reverse fault parallel to the Crackenback escarpment, was selected as correct by Cleary and others (1964). However, the results of recent geologic mapping (White and others, 1977) indicate that the Crackenback fault does not extend east of Lake Jindabyne. Moreover, the earthquake epicenters, as plotted, do not line up along the Crackenback fault. In fact, they are so widely scattered that selected groups could be attributed to the Jindabyne fault, the Berridale fault, and to a number of minor faults in the area between Berridale and Lake Eucumbene. Therefore, the alternative solution (involving a northeast or east-west trending, low-angle thrust fault) warrants further consideration. This solution would represent a plane subparallel to the Khancoban-Yellow Bog fault system in which bedrock has been thrust over Pleistocene(?) or more recent river gravels. This fault system appears to be related to the Long Plain fault, which is the most prominently visible structural feature on satellite imagery of the area (Figure 4-25).

Lake Talbingo

Lake Talbingo was outside the quadrangle of the original seismograph network, but a station was added in 1969, two years prior to filling. The additional station made it possible to locate events with magnitudes of 1 or less in this area. Prior to 1 May 1971, when filling began, seismic activity had been monitored for 13 years, including two years of nearby monitoring by the Talbingo station. The only earthquake recorded during that period was a minor one located about 19 km north of the dam site. Increased seismic activity commenced on 19 May 1971 with a small event and was followed by two more small earthquakes during May. The activity in June increased to 39 recorded events, of which four were of sufficient magnitude to be located (maximum magnitude recorded was 2.4). In July and August, an average of 20 locatable events per month occurred as reservoir filling continued. The rate of filling dropped sharply after August, and there was a correspondingly sharp decrease in the number of locatable events. However, the number of microearthquakes remained fairly constant, initially at several hundred per month (Timmel and Simpson, 1972). The locations of the larger earthquakes are shown on Figure 4-28).

Lake Blowering

Located immediately downstream from Talbingo Dam (Figure 4-25), the Blowering reservoir essentially has the same pre-impoundment seismic history as that described above for Lake Talbingo. During the 16-month period following impoundment, one microearthquake was located within 1 km of the reservoir.

FAULT INVESTIGATION

Lake Eucumbene Area

The geologic maps of the Lake Eucumbene area indicate that two faults, the Berridale wrench fault and the Jindabyne thrust fault, converge toward the Eucumbene Dam and that their joint continuation may extend beneath the reservoir (Figure 4-27). These features were examined during the field reconnaissance studies.

Berridale Wrench Fault - The Berridale wrench fault is a major northwest-trending feature described in the literature as exhibiting approximately 11 km of left-lateral offset and as showing evidence of vertical displacement (Lambert and White, 1965; White and others, 1977). The left-lateral displacement is believed to have occurred during the Devonian period, whereas the inferred vertical displacement apparently has offset basalts of Tertiary age. Clear evidence of this displacement was not observed during field reconnaissance conducted for the study. However, during aerial reconnaissance, the Berridale fault was recognizable on the basis of linear tonal changes in grassy areas, gross changes in vegetation, linear sidehill depressions, and aligned saddles (Figure 4-29).

Lambert and White (1965) state that the Berridale fault has produced linear scarps in alluvium southeast of Berridale, where the current reconnaissance started. They also show an aftershock sequence in the Berridale area as having a northwest-striking, left-lateral, stike-slip component. Thus, on the basis of stratigraphic, geomorphic, and seismic evidence, the Berridale fault was classified as active for the purposes of this study.

Jindabyne Thrust Fault - The Berridale fault appears to join the north-trending Jindabyne thrust fault immediately south of Eucumbene Dam (Figure 4-27). White and others (1976) state that the Jindabyne fault is marked by a prominent scarp over much of its mapped length (Figure 4-30), and they infer that this fault was a factor in impounding the postulated Pleistocene-aged Lake Jindabyne. White and others (1976) also

View is to the northeast. Arrows denote break in slope along Berridale fault.

BERRIDALE WRENCH FAULT
VICINITY OF LAKE EUCUMBENE
New South Wales, Australia

| | |
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| Project No. 14087A<br>**Woodward-Clyde Consultants** | Figure 4-29 |
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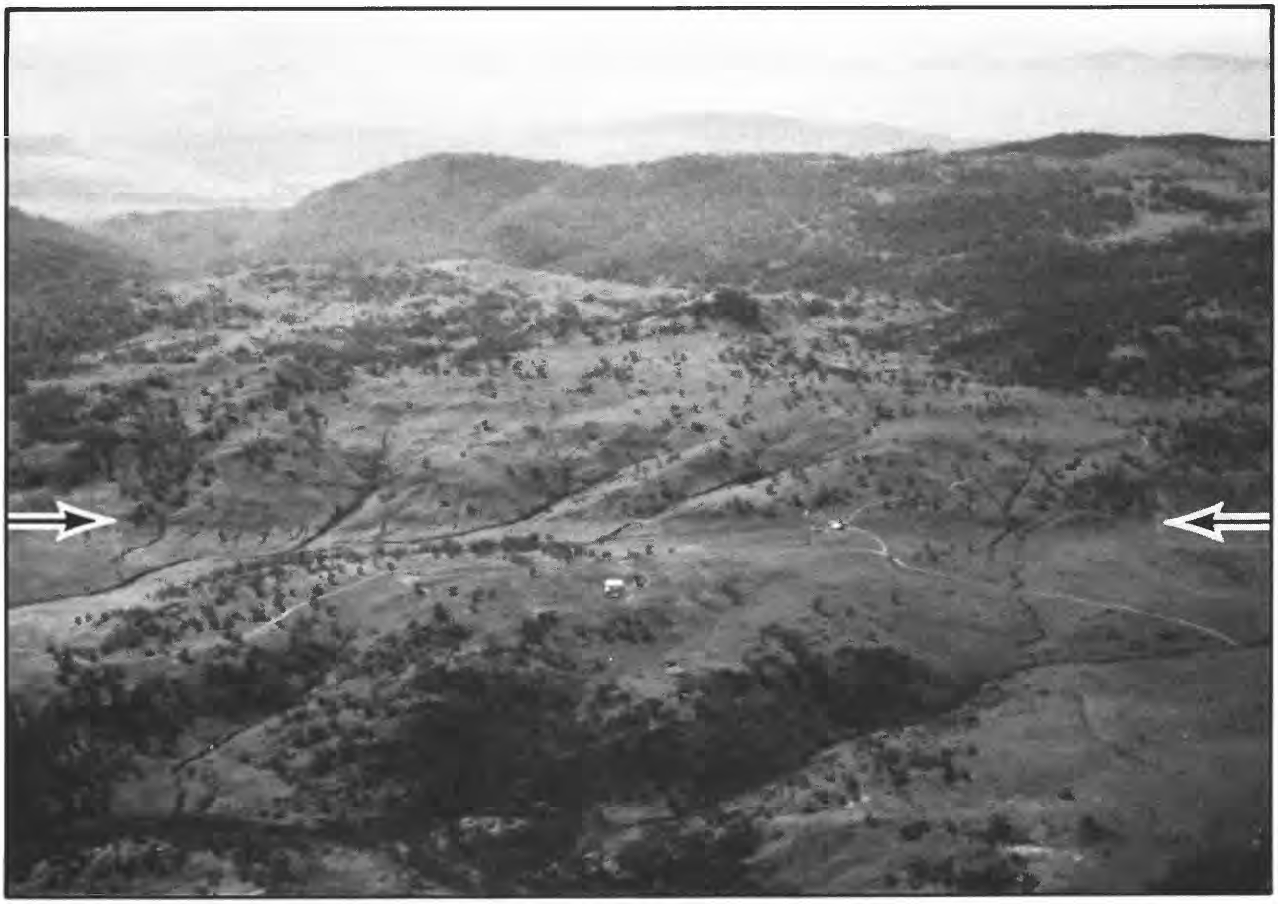


View is to the east, across Jindabyne fault near Hollins Crossing. The prominent topographic expression is shown.

| JINDABYNE THRUST FAULT<br>VICINITY OF LAKE EUCUMBENE<br>New South Wales, Australia | |
|---|---|
| Project No. 14087A<br>**Woodward-Clyde Consultants** | Figure 4-30<br>Page 184 |

point out that all faults mapped east and west of the Jindabyne fault are terminated by this fault. White and others (1976) identify a recent uplift along the Crackenback and Mowamba faults, both of which join or are cut off by the Jindabyne fault. The structural picture of the region, therefore, can best be explained as the Jindabyne fault being contemporaneous with the strike-slip faulting and having absorbed all the strike-slip displacement.

During the field reconnaissance for the current study, the mapped trace of the Jindabyne fault was observed as characterized by alignments of linearly grooved saddles, vegetational changes, and springs. For the purposes of this study, the Jindabyne fault is considered active, based upon the structural relationships.

Lake Talbingo and Blowering Areas

Long Plain Fault - Talbingo reservoir is located at the juncture of a "Y" that is formed by a splay in a major tectonic feature that could be construed to represent a fossil plate boundary (Figure 4-27). For the purposes of this report, this major tectonic line will be referred to as the Long Plain fault. (This nomenclature has been used previously by Maffi and Simpson, 1977.) This feature, which is shown on numerous geologic maps of the area (Pogson, 1972; Barnes and Herzberger, 1975; Brunker and others, 1970; and Degeling, 1977), extends southeastward into Victoria. In the project area, the Long Plain fault forms a dividing line between the Devonian-aged rock of the Kosciusko Plateau and Silurian- and Ordovician-aged rocks to the northwest (Figure 4-27). It is bordered by a broad zone of block faults and thrust faults. In the vicinity of Talbingo reservoir, this fault is characterized in part by the presence of a linearly extensive band of serpentinite and ultrabasic igneous rocks in a series of outcrops that are prominently visible from the air.

Individual faults within or evidently closely related to this system have been inferred to have been reactivated in post-mid-Tertiary time. K. R. Sharp of the Snowy Mountains Engineering Corporation (personal communication, 1978) cites as evidence of this reactivation thrusting of bedrock over gravel along the Khancoban-Yellow Bog fault in the Khancoban area (Figure 4-31) and a 125-m scarp that is across the Long Plain fault in 20 million-year-old basalt near Tumut Pond. He also cites extraordinarily deep accumulations of post-mid-Tertiary warping, or "tectonic ponding," that could be construed to suggest activity of faults along the major tectonic line. Gill and Sharp (1956) cite data establishing the age of the Kiandra deposits as Tertiary. However, the gravels at Khancoban are in a currently active river valley and may represent Pleistocene or younger deposition. Unfortunately, the exposure exhibiting bedrock thrust over gravel has become overgrown since the visit by Sharp, and it could not be located during the recent field studies. One road cut exhibiting evidence suggestive of thrusting of decomposed granite over gravels was observed near Khancoban, but the evidence was equivocable.

An unmapped feature that may represent an active branch of the Long Plain fault is present near the shore of Blowering reservoir. This feature is identifiable from the air on the basis of an alignment of linear depressions, saddles, and blocked or offset drainages. For the purposes of this study, the geomorphic and structural relationships along faults associated with the Long Plain fault have been considered to be evidence of fault activity during the active tectonic regime.

Bedrock thrust over gravel, vicinity of Khancoban. Photograph courtesy of Mr. Kenneth R. Sharp, Snowy Mountain Engineering Corporation, Cooma, N.S.W., Australia

| KHANCOBAN-YELLOW BOG FAULT<br>SNOWY MOUNTAINS<br>New South Wales, Australia | |
|---|---|
| Project No. 14087A<br>**Woodward-Clyde Consultants** | Figure 4–31<br>Page 187 |

# CONCLUSIONS

## Lake Eucumbene

The most extensive faults to which earthquakes in the Lake Eucumbene area could be attributed are the Jindabyne and Berridale faults. The mapped geologic relationships and geomorphic features observed during the aerial and field reconnaissance for the present study indicate that these faults are relatively young or have had a history of activity that has continued into late Cenozoic time.

## Lakes Talbingo and Blowering

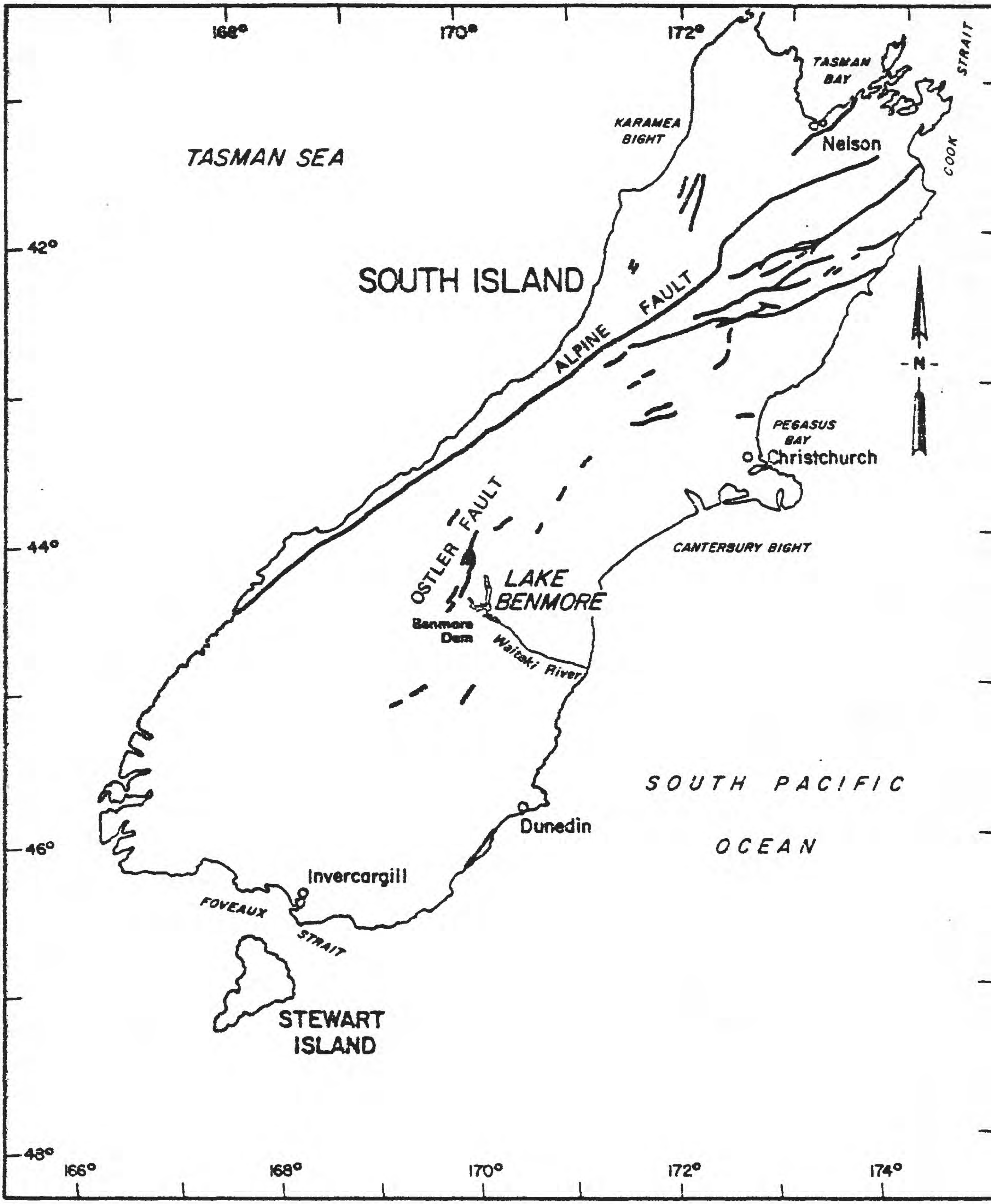
The seismic activity associated with the filling of Lake Talbingo may be related to faults associated with the Long Plain fault in the Lakes Talbingo and Blowering area. On the basis of the geomorphic evidence along these faults, and evidence of thrusting of bedrock over Pleistocene(?) or younger gravel along the Khancoban-Yellow Bog fault in the Khancoban area, these faults have been considered active.

## 4.5 LAKE BENMORE, NEW ZEALAND

### INTRODUCTION

Benmore Dam impounds the Waitaki River and its six tributaries, forming Lake Benmore, the largest man-made lake in New Zealand (Figures 4-32, 4-33, and 4-34; Table 4-3). Lake Benmore, located in the central portion of South Island (Figure 4-32), is a key feature of the Upper and Middle Waitaki Power Scheme. Two other large reservoirs, Lakes Pukaki and Ohau, are located on the headwaters of the Waitaki River within 25 km of Lake Benmore, and a third, Lake Tekapo, lies within 50 km of Lake Benmore (Figure 4-33). These three reservoirs were natural lakes that have been modified to serve as reservoirs.

After impoundment of Lake Benmore in 1964, seismicity within 80 km of Benmore Dam (including the area of the three subsidiary reservoirs) increased three to six times over the pre-impoundment seismicity (Adams, 1974). Because of this increase in seismicity upon the impoundment of Lake Benmore, a preliminary study was undertaken to evaluate if active faults are present within the influence of Lake Benmore. For the purposes of this study, a fault is defined as active if displacement has occurred during that area's present stress regime. For the Lake Benmore area, this would be Late Cenozoic time.




EXPLANATION:

Active Fault

0 20 40 60 80 100 km

LOCATION MAP OF LAKE BENMORE

South Island, New Zealand

| Project No. 14087A | Figure 4–32 |
|---|---|
| Woodward-Clyde Consultants | Page 190 |


| PROJECT AREA<br>LAKE BENMORE<br>South Island, New Zealand | |
|---|---|
| Project No. 14087A<br>**Woodward-Clyde Consultants** | Figure 4–33<br>Page 191 |

View is to the north, across Lake Benmore: Benmore Dam is in foreground and Lake Pukaki in distance.

| BENMORE DAM AND LAKE BENMORE South Island, New Zealand | |
| --- | --- |
| Project No. 14087A **Woodward-Clyde Consultants** | Figure 4–34 |
| |  |

TABLE 4-3
DESCRIPTION OF THE
BENMORE DAM AND RESERVOIR

| | |
|---|---|
| Type | Earthfill |
| Structural Height | 110 m |
| Hydraulic Head | 95 m |
| Crest Length | 820 m |
| Reservoir Area | $2.04 \times 10^9$ $m^2$ |
| Gross Storage | $2.2 \times 10^8$ $m^3$ |
| Date Completed | November 1964 (filling started) |
| Foundation Rock | graywacke and argillite |

This investigation incorporates the results of studies of seismicity and faulting in the vicinity of Lake Benmore by personnel of the New Zealand Geological Survey in New Zealand. The studies for this report were selective and limited in scope; it was not intended that all faults in the area be identified and evaluated. For the purposes of this study, it was assumed that any Late Cenozoic faults within the hydrologic regime of the reservoir may be influenced by impoundment of the reservoir.

The scope of this study's work consisted of the following:

- gathering and reviewing published and unpublished data and maps;

- questioning scientists resident in, or familiar with, the study area;

- making aerial reconnaissance flights over the study area;

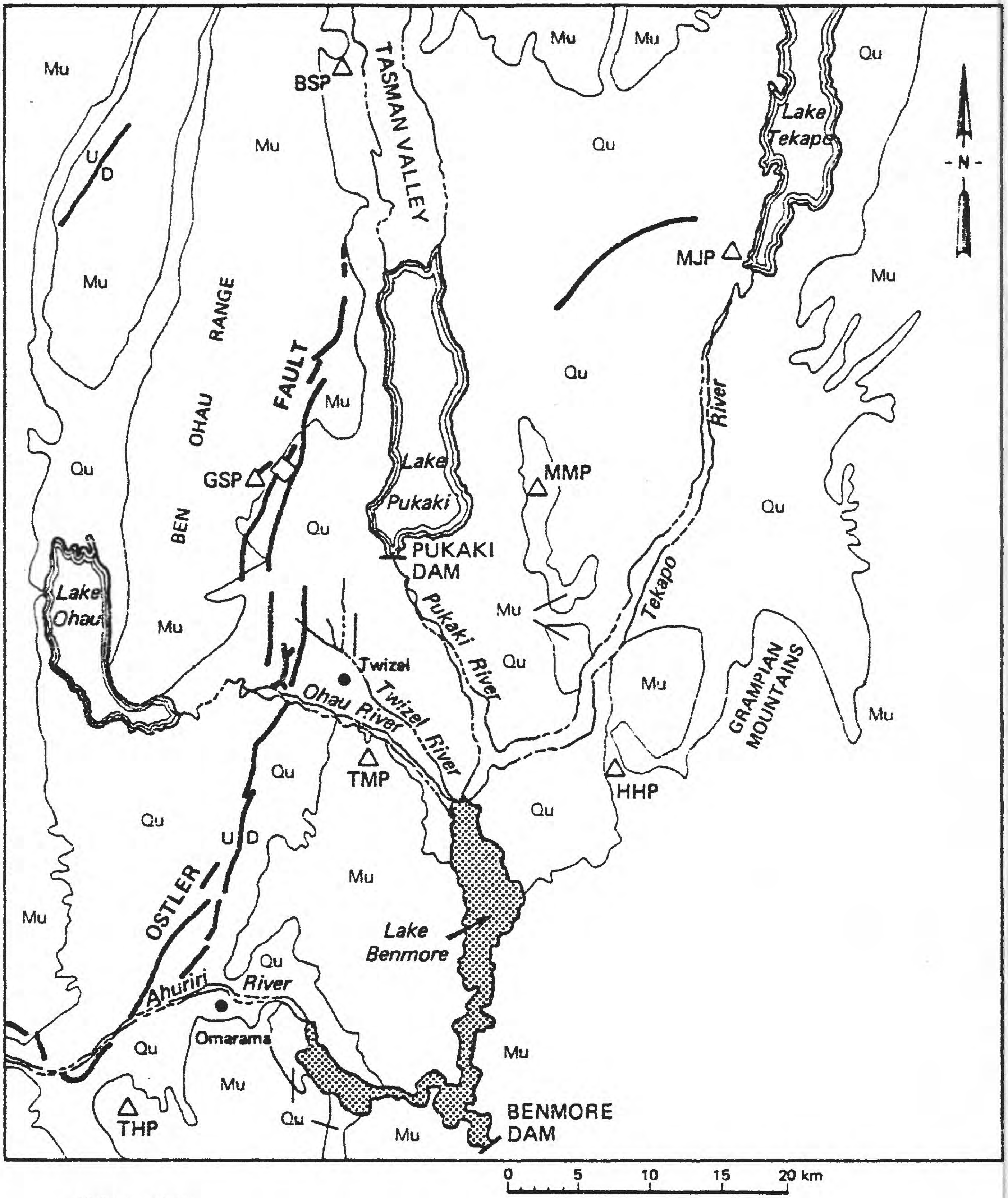
- performing ground reconnaissance at selected locations in the study area;
- examining LANDSAT imagery and aerial photographs of portions of the study area; and
- collecting and evaluating seismic data.

The initial part of the data gathering effort was performed in our offices in San Francisco, using Woodward-Clyde Consultants' library facilities and files. Additional data and maps were obtained from government agencies in Lower Hutt, Wellington, and Twizel, New Zealand. Selected references also were reviewed in the library of the New Zealand Geological Survey, and complete seismologic data were obtained from the Seismological Observatory in Wellington. A reservoir filling history was requested from the Ministry of Works and Development, and aerial photographs were purchased from the Department of Lands and Survey.

## REGIONAL GEOLOGIC SETTING

Benmore Dam and Reservoir area is a complexly faulted basin and has a range topography in which regional uplift is believed to have continued into the Pleistocene period (Shaw and Stevens, 1966). However, much of the faulting in the region probably occurred during Tertiary time (McKellar and others, 1967).

The only recognized major historically active fault in the region is the Alpine fault, which is located 32 km northwest of Benmore Dam (Figure 4-32). However, evidence of Holocene activity along segments of other faults of potentially significant extent has been reported (Gair, 1967). The Ostler fault zone (Figure 4-35), which is described in detail later in this report, is an example.




**EXPLANATION:**

Geologic contact

U/D Active fault trace; relative movement indicated

△ Microseismic station

◇ Epicenter of December 1978 M 4.6 earthquake

Qu Undifferentiated Quaternary deposits, principally alluvium, till, and glacial outwash

Mu Undifferentiated Mesozoic bedrock, principally graywacke and argillite

Source: Gair (1967), Mutch (1963), and MacFarlane (1979)

REPORTED ACTIVE FAULT TRACES
VICINITY OF LAKE BENMORE
South Island, New Zealand

Project No. 14087A

Woodward-Clyde Consultants

Figure 4–35

The available geologic maps (Mutch, 1963; Gair, 1967) indicate that Benmore Dam and much of Lake Benmore are located on graywacke and argillite of Silurian age (Figure 4-35). A relatively minor portion of the reservoir is mapped as being located on similar materials of Permian age. The northern end of the reservoir is mapped as being located on Pleistocene alluvium and glacial till. The Pleistocene deposits are believed to reach a thickness in excess of 1518 m upstream from the reservoir (Mansergh and Read, 1973). A total of seven Pleistocene units have been recognized in the project area. In general, these consist of glacial till, outwash gravels, interglacial deposits, and alluvium.

REGIONAL SEISMIC ACTIVITY

Lake Benmore is located in an area of relatively low seismicity. During the pre-impoundment period (1955 to 1964), only 7 earthquakes of magnitude 4 or greater were reported to have been located within 80 km of Benmore Dam (Adams, 1974). During the post-impoundment period (1965-1972), a total of 29 earthquakes occurred within 80 km of the dam. Two of these events had a magnitude of 5, whereas the highest magnitude recorded between 1955 and 1964 was 4.4.

Earthquakes within radii of 40 and 60 km also show a statistical difference between pre-impoundment and post-impoundment seismicity: the number of earthquakes having magnitudes equal to or greater than 4 has increased by factors of 3.1 and 6.5 since impoundment of Lake Benmore. The factor of 3.1 applies to earthquakes within 60 km of the dam, and the factor of 6.5 relates to earthquakes within 40 km of the dam. In considering distance and azimuth together, Adams (1974) found that 50 percent of the earthquakes examined occurred in 10 percent of the area considered (upstream of the dam and within 40 km of it). However, reservoir impoundment in

December 1964 was not followed immediately by an increase in earthquake activity; the first magnitude 5 event did not occur until July 1966. The second magnitude 5 event occurred nearly 5 years later, in April 1971. Neither event could be correlated with changes in reservoir water level, as it has remained essentially constant since the intitial filling (Adams, 1974).

Woodward-Clyde Consultants reexamined the data base and incorporated additional data with that studied by Adams (1974). The locations of these events are shown on Figure 4-36.

The first magnitude 5 earthquake to occur following impoundment of Lake Benmore was plotted near the Otemata fault, about 17 km west of the dam (Adams, 1974). This fault is not mapped as a Quaternary feature (New Zealand Geological Survey, 1973). The second magnitude 5 earthquake to occur following impoundment was located approximately 14 km north of the dam, adjacent to the western shore of the reservoir and away from any known fault. This event suggests the possibility of the existence of a fault along the margin of the mountains where they join the Mackenzie Basin (Adams, 1974). A number of smaller events fall within a northeast-trending alignment, extending roughly along the mountain front (Adams, 1974). No surface evidence of the existence of a fault in this location has been reported.

Several post-impoundment earthquakes have been plotted within or adjacent to the Ostler fault zone. This zone, including its probably northern extension, is the only relatively well-documented Quaternary fault in the project area (Mansergh and Read, 1973). The largest and most recent earthquake that appears to be related to this fault zone was a magnitude 4.6 event that occurred on 17 December 1978. The epicenter was located on the northern extension of the Ostler fault zone, west of Lake Pukaki (Figure 4-36). The available data suggest a focal depth of approximately 4 km. Personnel of the New Zealand Geological Survey conducted a detailed examination of known fault traces in the epicentral area but found no evidence of surface fault rupture or other ground damage (Don Macfarlane, personal communication, December 1978). The pattern of first motions is reported to be consistent with sinistral strike-slip movement on a north-south fault. This interpretation is supported by the strong motion data, which showed most of the seismic energy to be in a north-south direction and of anomalously low amplitude in the epicentral area (Calhaem, 1978).

The seismic energy released in the December 1978 earthquake is reported to be more than 600 times the total energy released during the 3 1/2 year period immediately preceding that event (Calhaem, 1978). This earthquake occurred during rapid filling of Lake Pukaki. Lake Pukaki is a natural lake, the capacity of which has twice been increased by construction of dams designed to increase the power production capacity of the Upper and Middle Waitaki Power Scheme. The new Pukaki High Dam, which was completed in 1978, ultimately will create a reservoir with a maximum depth of 108 m and a storage capacity of more than $10 \times 10^9$ $m^3$. As a precaution against further seismic activity, the rate of filling of this reservoir was reduced by 50 percent following the 17 December 1978 earthquake (Calhaem, 1978).

During 1973, the New Zealand Institute of Geophysics recorded an 18-day survey of microearthquakes in the project area. This was done to establish in more detail the locations and mechanisms of earthquakes in the Benmore area and to ascertain the base level of seismicity prior to raising the level of Lake Pukaki. This portable network was capable of detecting


HISTORICAL MACROSEISMICITY
VICINITY OF LAKE BENMORE
South Island, New Zealand

Woodward-Clyde Consultants

Figure 4-38

and locating seismic events with magnitudes less than 0, as ompared to a higher limit of magnitude for the pre-existing fixed network. Two linear trends were identified from plotting located events. One extended in a northerly direction, from the vicinity of the town of Omarama to the Ben Ohau Range (Figure 4-35); this trend corresponds quite closely with the mapped position of the Ostler fault zone. The second trend extended in a northeasterly direction, from the northern arm of Lake Benmore along the edge of the Grampian Mountains (Figure 4-35). The most active areas along these two trends correspond closely to the postions of the two magnitude 5 earthqakes that had been recorded previously (Adams and others, 1974).

Although the main trend of epicenters appears to correspond closely with the Ostler fault zone, the calculated mechanisms indicate that the fault planes of individual earthquakes strike approximately 45 degrees from this alignment. This strike is consistent with that of the trend noted along the base of the Grampian Mountains. Both the computer microearthquake mechanisms and the geologically observed faulting are consistent with east-west regional compression (Adams and others, 1974).

FAULT INVESTIGATION

A small-scale map of active faults in New Zealand (Lensen, 1965) indicates the presence of a number of relatively minor active faults in the region surrounding Lake Benmore. The most extensive of these, and the one closest to the reservoir, is the Ostler fault zone (Figure 4-35). It is located approximately 10 km west of the western arm of Lake Benmore and approximately 18 km west of the center of the main body of the reservoir. It extends for at least 56 km, from the Ahuriri River on the south to the headwaters of the Twizel

River on the north (Gair, 1967). From there, the mapped relationships suggest that the fault may join a less recent fault system that extends northward into the Southern Alps, where it dies out in a fold. The axis of this fold is located subparallel to, and approximately 21 km southeast of, the Alpine fault, which is the most significant tectonic feature in New Zealand. The northern extension of the Ostler fault is approximately 40 km long and has been mapped as being active during Pleistocene, but not Holocene, time (New Zealand Geological Survey, 1973).

The Ostler fault zone, due to its proximity to a power plant site, has been studied in some detail by the New Zealand Geological Survey. These studies have included exploratory trenching and profiling. As described by Mansergh and Read (1973), the Ostler fault zone is made up of a series of individual traces within a band that locally may be as much as 5500 m wide (Figure 4-37). The most prominent of these traces in the study area are, from west to east, the Haybarn fault, the "Y" fault, and the Ruataniwha fault (Figures 4-38 and 4-39). Mansergh and Read (1973) state that the individual fault blocks within the zone generally are downthrown to the east and suggest that the locus of movement has migrated to the east with time. Tilted surfaces and folds are present within the zone.

The New Zealand Geological Survey (Mansergh and Read, 1973) has recognized two and possibly three areas where displacement of 1.8 to 2.7 m in alluvium occurred along the Ostler fault zone. Moreover, a total maximum displacement of 20 m was measured in 14,000-year-old glacial deposits, and 21 m in 16,000-year-old glacial deposits. On this basis, the average rate of fault movement has been estimated to be 14 m per thousand years. However, based upon maximum and minimum recognized individual episodes of offset of 6.1 to 6.7 m and

View is to the south, along the Ostler fault zone. Arrows indicate variations in scarp height in Quaternary units of different ages.

| OSTLER FAULT ZONE<br>VICINITY OF LAKE BENMORE<br>South Island, New Zealand | |
|---|---|
| Project No. 14087A<br>**Woodward-Clyde Consultants** | Figure 4–37<br>Page 202 |

View is to the north, along the Haybarn trace of the Ostler fault zone. The left-lateral drainage offset can be seen near excavation scar.

| HAYBARN FAULT<br>VICINITY OF LAKE BENMORE<br>South Island, New Zealand | |
|---|---|
| Project No. 14087A<br>**Woodward-Clyde Consultants** | Figure 4–38<br> |

View is to the north, along the Haybarn trace of the Ostler fault zone. The degraded fault scrap is visible. Left-lateral offset of small drainage crosses slope adjacent to geologist (in center of photograph).

| HAYBARN FAULT<br>VICINITY OF LAKE BENMORE<br>South Island, New Zealand | |
|---|---|
| Project No. 14087A<br>**Woodward-Clyde Consultants** | Figure 4–39<br>Page 204 |

1.8 m, respectively, the recurrence interval of episodes of surface fault rupture has been estimated as 2800 years (Mansergh and Read, 1973).

The studies performed by the New Zealand Geological Survey indicate that the Ostler fault zone is the result of thrusting, with the west side moving up. Lake Benmore is located on the lower plate of the thrust, approximately 15 km to the east of the zone. An alignment of earthquake epicenters, all recorded since the filling of the reservoir in 1964, have been mapped within or immediately west of the zone. This seismicity has been described in detail by Adams (1973) and has been summarized previously in this report.

The Ostler fault zone was examined on the ground, from the air, and on aerial photographs. The three fault traces mapped in the Twizel area by the New Zealand Geological Survey (Mansergh and Read, 1973) are prominently visible in the aerial photographs. On these photographs and from the air, the Ostler fault zone is identifiable principally on the basis of aligned scarps. Some of the scarps are substantially higher than others and may represent erosional remnants. The ground surfaces on the western side of both the high scarps and the low scarps have, in general, a gentle westward tilt. This westward tilt also was observed in bedding planes exposed in road cuts and stream banks. Landslides were noted on a number of the higher scarps; however, no landslides were observed at locations away from the fault zone.

Some of the scarps, particularly the lower ones, are quite irregular and tend to follow the trace inferred by the New Zealand Geological Survey (Mansergh and Read, 1973). However, in gross aspect, the zone is quite linear. This is particularly true west of Lake Pukaki where the probable northern extension of the Ostler fault zone appears essentially as a straight line marking a break in slope along a precipitous mountain front. This pronounced linear trend, particularly when considered with the reported alignment of earthquake epicenters along the surface trace of the Ostler fault zone, strongly suggests that the major portion of the fault plane is nearly vertical, although it may "roll over" somewhat where it passes through deep alluvial and glacial deposits.

For the present study, the surface reconnaissance of the Haybarn trace of the Ostler fault zone found pronounced evidence of left-lateal ofset of small drainages. The amount of left-lateral offset appears to be substantially greater than the vertical displacement. This evidence of a strong left-lateral strike-slip offset supports the previously described evidence pointing to the existence of a nearly vertical fault plane. It also supports the focal mechanism interpretation for the December 1978 earthquake. The strongest geomorphic expression suggestive of recurrent offset was seen on the Ruataniwha fault trace, immediately north of the Ohau River. At this location, two fault traces appear to be closely intertwined, producing a scarp-like feature. The results of the present study and of evaluation of the mechanism of the December 1978 earthquake indicate that the Ostler fault zone, which previously had been considered to be a westward-dipping thrust fault, is a nearly vertical, left-lateral strike-slip fault along its eastern portion.

CONCLUSIONS

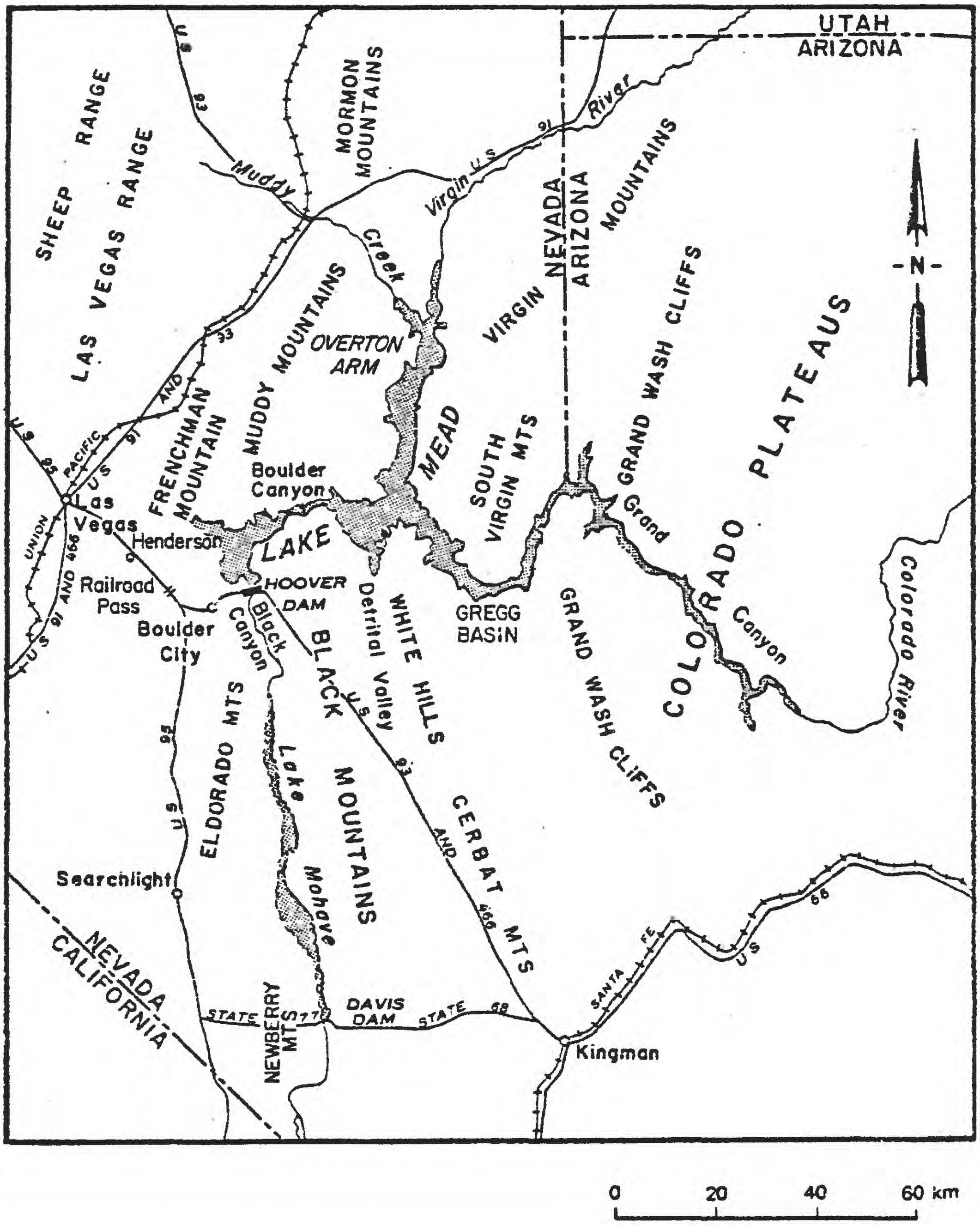
The available data demonstrate a pronounced and significant increase in local seismicity following impoundment of Lake Benmore. Some of the earthquakes that have occurred since impoundment of Lake Benmore and that have been suspected of being induced by it have been located along or closely adjacent to the Ostler fault zone. (The center of Lake Pukaki is located about 8 km to the east of the Ostler fault zone.) The Ostler fault zone contains faults that have had surface displacement during the current tectonic regime, as expressed by scarps in Pleistocene deposits and on Holocene surfaces. The most recent earthquake on this fault, which occurred on 17 December 1978, may have been triggered by rapid filling of Lake Pukaki.

## 4.6 HOOVER DAM AND LAKE MEAD, UNITED STATES

### INTRODUCTION

Lake Mead, impounded by construction of Hoover Dam across the Colorado River at the Arizona-Nevada border in Black Canyon (Figure 4-40), is one of several large, deep reservoirs that have been responsible for inducing or triggering an increased level of local seismicity. This report describes the results of preliminary studies undertaken to evaluate the possibility that the reservoir is underlain or closely bordered by one or more active faults. For the purposes of this study, an active fault is one on which displacement has occurred during that area's present stress regime; for the Lake Mead area, this would be Late Cenozoic time.

In part, these studies are an extension of earlier work by others who had correlated some of the recorded seismic events with mapped geologic structures. However, the present studies were selective and limited in scope; it was not intended that all faults in the area be evaluated. For the present study, a brief review of pertinent available data was made. The locations of selected faults and seismically active areas described in the literature were marked on topographic maps to facilitate the field studies. Selected areas were examined for regional trends on aerial photographs (scale 1:80,000) provided by the U.S. Geological Survey. Areas and linear trends of potential interest and significance were then viewed during aerial reconnaissance flights. Selected areas accessible by foot or four-wheel drive vehicle were examined during the ground reconnaissance. Extensive 35 mm photo coverage of the study areas was taken for reference during data evaluation.




Source: After Longwell (1963)

LOCATION MAP
HOOVER DAM AND LAKE MEAD
Arizona and Nevada, United States

Project No. 14087A
**Woodward-Clyde Consultants**

Figure 4–40

REGIONAL GEOLOGIC SETTING

Hoover Dam impounds the Colorado River in northwestern Arizona and southeastern Nevada forming Lake Mead; the maximum depth of the reservoir is approximately 166 m. The dam and reservoir are located in the Basin and Range physiographic province, which is characterized by regional extension, resultant normal faulting, and horst and graben structures (Figure 4-41). The dam itself is underlain by Tertiary volcanic deposits of basalt and andesite. The reservoir also is underlain by Tertiary volcanic deposits and by Tertiary conglomerate, sandstone, clay, salt, and gypsum deposits, Precambrian metamorphic units, and Precambrian granite (Longwell, 1936; Rogers and Lee, 1976).

Faulting is prevalant in the vicinity of the dam site and in the reservoir area. The faults are late Tertiary in age and are primarily strike-slip faults that trend northeast and normal faults that trend north (Anderson and Laney, 1975). The two major strike-slip faults, the Las Vegas shear and the Hamblin Bay fault, are estimated by Rogers and Lee (1976) to be approximately 64 to 80 km long. Cumulative displacements on the normal faults measure approximately 1800 m, according to the same authors. A large number of faults of lesser extent have been mapped (Figure 4-41).

REGIONAL SEISMIC ACTIVITY

Prior to reservoir impoundment in 1936, no historical earthquakes had been reported near the reservoir, and the area was considered to be one of extremely low seismicity. Seismic activity, which commenced subsequent to reservoir impoundment, was clustered within 25 km of the reservoir. During an 8-year monitoring period beginning in 1937, epicenters were centered primarily along existing faults at focal depths of less than


NOTE:
Lineaments indicated by arrows.

| LIME RIDGE AREA<br>HOOVER DAM AND LAKE MEAD<br>Arizona and Nevada, United States | |
|---|---|
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| | Page 217 |